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Andrey G. Kalinichev

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1 MOLECULAR STRUCTURE AND DYNAMICS OF NANO-CONFINED WATER:
2 COMPUTER SIMULATIONS OF AQUEOUS SPECIES IN CLAY, CEMENT, AND POLYMER MEMBRANES

3 Andrey G. Kalinichev^{1,2}

4 ¹ Laboratoire Subatech, UMR 6457, Ecole des Mines de Nantes, 44307 Nantes, FRANCE

5 ² Department of Chemistry, Michigan State University, East Lansing, MI 48824, USA

6 Abstract. Molecular-level knowledge of the thermodynamic, structural, and transport properties of water confined by
7 interfaces and nanopores of various materials is crucial for quantitative understanding and prediction of many natural
8 and technological processes, including carbon sequestration, water desalination, nuclear waste storage, cement
9 chemistry, fuel cell technology, etc. Computational molecular modeling is capable to significantly complement the
10 experimental investigations of such systems by providing invaluable atomic-scale information leading to improved
11 understanding of the specific effects of the substrate structure and composition on the structure, dynamics and reactivity
12 of interfacial and nano-confined aqueous solutions. This paper offers a brief overview of recent efforts to quantify some
13 of these effects for individual H₂O molecules and hydrated ions confined at the interfaces and in nanopores of several
14 typical hydrophilic and hydrophobic materials. The first molecular layer of aqueous solution at all substrates is often
15 highly ordered, indicating reduced translational and orientational mobility of the H₂O molecules. This ordering cannot
16 be simply described as “ice-like”, but rather resembles the behavior of supercooled water or amorphous ice, although
17 with very significant substrate-specific variations.

18 1. INTRODUCTION

19 Most chemical reactions near the Earth’s surface involve fluid (aqueous) phases and take place at fluid-solid
20 interfaces or in confined spaces of mineral interlayers and nanopores. These reactions affect many important natural
21 processes, including mineral weathering, adsorption or release of environmental contaminants in soil, drinking water
22 quality, formation and behavior of ice nano-crystals and hydrated mineral nano-particles in the atmosphere, the fate of
23 CO₂ in geologic carbon sequestration. The same interfacial phenomena are also at the core of many important
24 technological processes related to water purification and desalination membranes, fuel cells, production and utilization
25 of cement, etc. All these phenomena are inherently multi-scale in time and space, and addressing the scientific and
26 technological problems they pose is multidisciplinary. Molecular-level knowledge of the chemistry and physics of
27 interfacial and confined aqueous solutions interacting with mineral surfaces and other substrates is essential to their
28 understanding. Important questions center on how water and dissolved species behave at the surfaces and in nano-
29 confinement by different inorganic and organic substrates.

30 Current understanding of hydrated mineral-water interfaces is far from complete. Molecular-scale information
31 is difficult to obtain experimentally due to the nature of the materials (e.g., interfacial and bulk structural disorder, the
32 presence of molecular scale dynamical disorder with characteristic time scales that span many orders of magnitude, the
33 physical limitations on the samples that can be effectively examined, and the difficulty of experiments under extreme
34 conditions). Mineral surfaces themselves vary by the arrangements and reactivity of individual surface functional
35 groups and the specific adsorption of inorganic and organic ions. Even the utility of macroscopic thermodynamic
36 concepts, such as the dielectric constant and electric double-layer, remain open and are important questions driving the
37 current research [1-3]. Modern experimental probes, including synchrotron light sources, high-resolution neutron
38 scattering, time-resolved spectroscopy and diffraction, scanning force microscopy, and multi-dimensional NMR
39 spectroscopy, are providing unprecedented, atomic-level and surface-specific information about mineral-fluid systems
40 [3-5]. However, it is often difficult to interpret these experimental observations without parallel theoretical and
41 computational studies [6-7]. In addition, the interaction with experimentalists is imperative for verifying theoretical
42 models and guiding larger scale simulation efforts. Connecting computational and experimental results is crucial in
43 addressing these questions and in bridging the gap between the atomic- and nanometer-scales and the micron- and
44 larger scales that is essential for environmental, geochemical and energy applications of practical importance [8].

45 The methods of atomistic computer simulations [9], often coupled to surface-specific experimental results, are
46 already mature enough to provide otherwise unobtainable perspective into the structure, dynamics and energetics of
47 fluid-mineral systems [10]. Computational approaches range from traditional quantum chemical methods to classical
48 molecular dynamics (MD) and Monte Carlo (MC) simulations using semi-empirical interaction potentials to emerging
49 methods including quantum molecular dynamics, metadynamics, and reactive force fields for molecular-scale modeling
50 [11-35]. Critical questions currently revolve around the effect of solid surfaces on the properties of near-surface and

51 nano-confined fluids, structural and chemical modification of the surface by the fluid, the pathways of chemical
52 reactions at surfaces and in fluids, the connections between these reactions and fluid structure and dynamics, and the
53 consequent effects on the pathways of dissolution and precipitation. This paper provided a brief overview of our
54 group's recent efforts to quantify some of these properties and processes at the interfaces and in nano-confinement with
55 several typical hydrophilic and hydrophobic materials by classical MD simulations using the CLAYFF force field [12].

56 2. COMPUTATIONAL MODELS AND METHODS

57 We compare the molecular structure of several qualitatively different substrate-water interfaces. Crystals of
58 metal hydroxides, like **brucite**, $\text{Mg}(\text{OH})_2$, or **portlandite**, $\text{Ca}(\text{OH})_2$, can be naturally cleaved parallel to their (001)
59 crystallographic plane, resulting in an electrostatically neutral surface fully covered by OH groups oriented away from
60 the surface. They can easily form hydrogen bonds with interfacial H_2O molecules. Therefore, we will use the surface of
61 portlandite as a reference model of a hydrophilic interface [11,20]. A similar hydrophilic surface can be created by
62 cleaving the crystal structure of **quartz**, SiO_2 , parallel to the (001) plane. In the presence of water and at relatively low
63 pH, this surface will also be fully covered by OH groups [36]. In the structure of **hydrocalumite**,
64 $\text{Ca}_2\text{Al}(\text{OH})_6\text{Cl}\cdot 2\text{H}_2\text{O}$, 1/3 of the divalent Ca cations are isomorphically substituted by trivalent Al, resulting in a
65 positive structural charge compensated by the presence of Cl^- or some other anions in the interlayer space and on the
66 surfaces of this material, an anionic clay [11,20]. Similar to its neutral portlandite analog, its cleaved (001) surface is
67 also fully covered by OH groups and can be considered hydrophilic.

68 The structure of **talca**, $\text{Mg}_3\text{Si}_4\text{O}_{10}(\text{OH})_2$, consists of so-called TOT layer (an octahedral layer of $[\text{MgO}_4(\text{OH})_2]$
69 sandwiched between two tetrahedral layers of $[\text{SiO}_4]$) stacked along the crystallographic Z direction [18,21,27]. All
70 tetrahedral sites are occupied by Si forming $[\text{SiO}_4]$ structural units, and all octahedral sites are occupied by Mg forming
71 $[\text{MgO}_4(\text{OH})_2]$ units. The TOT layers are, thus, electrostatically neutral and interact with each other only by weak van
72 der Waals forces, reflecting the hydrophobic nature of the basal siloxane Si-O-Si (001) surface. The TOT layered
73 structure is typical for many clay minerals, which can develop various degrees of negative structural charge due to the
74 isomorphous substitutions possible for both tetrahedral (e.g., Al for Si) and octahedral (e.g., Li for Mg) layers. The
75 emergence of negative layer charge is balanced by the presence of charge-balancing cations in the interlayer space and
76 on the surfaces of these minerals and makes them more hydrophilic. Their interaction with water has already been
77 extensively investigated [12,25,26,29,30,32,37] and we focus here only on the behavior of neutral talc, as a reference
78 hydrophobic surface.

79 It is known experimentally (e.g., [38]) that mineral surfaces rarely exist in nature in the atomically smooth
80 ideal form as the ones listed above. To illustrate the effects of nanoscale surface roughness of the substrate on the
81 structural and dynamic properties of the interfacial aqueous solution, the surface of **tobermorite**, $\text{Ca}_5\text{Si}_6\text{O}_{16}(\text{OH})_2$, is
82 taken here for comparison. Tobermorite is also known as one of the so-called C-S-H (calcium silicate hydrate) phases in
83 cement chemistry and its surface properties control behavior of confined water in the nanopores of cement and concrete
84 [11,20].

85 The disordered interfaces of water filtration membranes are also typically characterized by the surface
86 roughness on the nanoscale and above. In this case, the diversity of the surface adsorption sites can be very large and it
87 is not always possible to identify and investigate all of them individually at a truly atomistic level. Nevertheless,
88 computer simulations can still provide invaluable information about the molecular mechanisms of the processes of
89 membrane filtration. The effects of various metal cations and their interactions with dissolved natural organic matter on
90 the process of bio-fouling of **polyethersulfone (PES) desalination membrane** is taken here to illustrate the case [39].

91 All MD simulations described here were performed at ambient conditions in the *NVT* statistical ensemble
92 (constant number of atoms, constant volume, and temperature $T = 300$ K) after a pre-equilibration in the *NPT* ensemble
93 at $T = 300$ K and $P = 0.1$ MPa using standard MD algorithms [9]. The mineral surfaces were constructed by cleaving the
94 bulk crystal structures parallel to the (001) plane. The irregular polymer surface was created in the course of preliminary
95 MD runs of condensing a polyethersulfone polymer chain consisting of 60 monomer units [39]. Each complete
96 interfacial MD simulation cell consisted of a substrate slab about 2 to 3 nm thick, and a layer of liquid water about
97 3 to 7 nm thick placed in contact with them, as illustrated in Fig.1. The water was pre-equilibrated at 300 K and
98 0.1 MPa. Periodic boundary conditions [9] were then applied in all three dimensions to produce models of the interfaces
99 formed by infinite (in x - and y directions) flat substrate layers interspersed with layers of water. In all cases, the
100 thickness of the water layer and the size of the cell in the z -direction (normal to the surface) were sufficiently large to
101 effectively exclude direct interactions between two different substrate/water interfaces created due to the periodicity of
102 the system. The number of H_2O molecules in the layer was chosen to reproduce the density of bulk water under ambient
103 conditions (~ 1 g/cm³). The final x -, y -, and z -dimensions of the simulation supercells were approximately
104 $2 \text{ nm} \times 2.5 \text{ nm} \times 10 \text{ nm}$ with slight variations depending on the nature of the substrate.

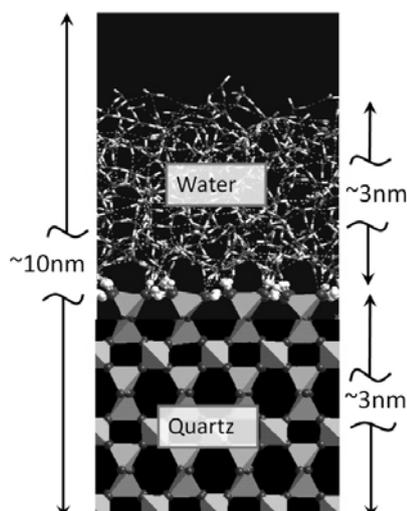


Figure 1: A schematic view of the molecular dynamics periodic simulation cell for the modeling of a substrate-water interface. Bulk quartz SiO_2 crystal is shown as tetrahedral (Si) and dark balls (O). White (H) and dark (O) balls represent the O-H groups of the fully hydroxylated surface. Approximately half of the surface O-H groups are bent to make hydrogen bonds among themselves, while the other half – donate H-bonds to the interfacial H_2O molecules, which are shown as white (H) and dark (O) cylinders above the surface. H-bonds are indicated as thin dashed lines.

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The MD time step was 1.0 fs, and each model system was allowed to equilibrate for at least 500 ps of MD before an equilibrium MD trajectory for each model was accumulated at 10 fs intervals during additional 500 to 1000 ps of MD simulation. The details of the MD simulation procedures and parameter settings for mineral-water systems are described in previous publications [11,12,18,20,27,37,39-41]. The structural analysis of the near-surface water films was undertaken using atomic density profiles in the direction perpendicular to the substrate surface and atomic density maps within defined slices of the water film parallel to the surface. These properties were all calculated by averaging over the last 500 ps equilibrium MD trajectory of each system. The position of the mineral surface ($z = 0$) was defined by the average position of the top-most surface oxygen atoms or of the corresponding OH groups, depending on the substrate.

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3. SIMULATED RESULTS AND DISCUSSION

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3.1. Interfacial water structure

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Fig. 2 shows the time-averaged atomic density contour map of O_w in the first molecular layer of water on the (001) surface of quartz. It is important to note, that unlike the surfaces of brucite [18] and portlandite [11], only about 50% of the surface hydroxyl groups on quartz are oriented away from the surface and easily donate H-bonds to the interfacial H_2O molecules. The other 50% of the surface OH groups are oriented parallel to the surface of quartz, form H-bonds to other surface hydroxyls, and accept H-bonds from H_2O molecules (Figs1,2). The surface ordering of the H_2O molecules is clearly observed in the time-averaged patterns of the O_w distribution. Surface water molecules are significantly immobilized by each accepting an H-bond from one surface OH group and simultaneously donating one H-bond to another neighboring surface OH group. A similar pattern of H_2O ordering is also observed at the neutral hydroxylated surfaces of brucite [18] and portlandite [11]. However, the charged hydroxylated surface of hydrocalumite exhibits a completely different behavior where no H-bonds donation is possible for the interfacial water molecules [11].

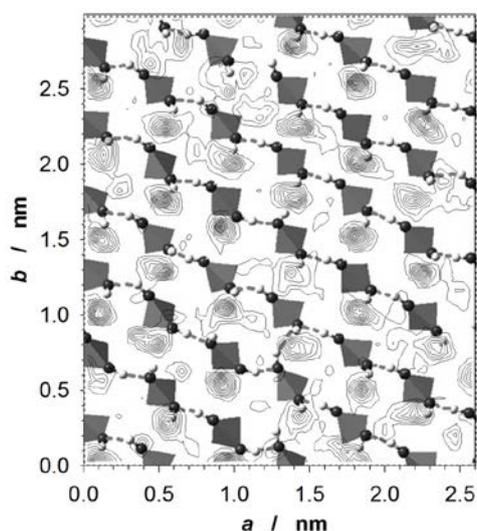
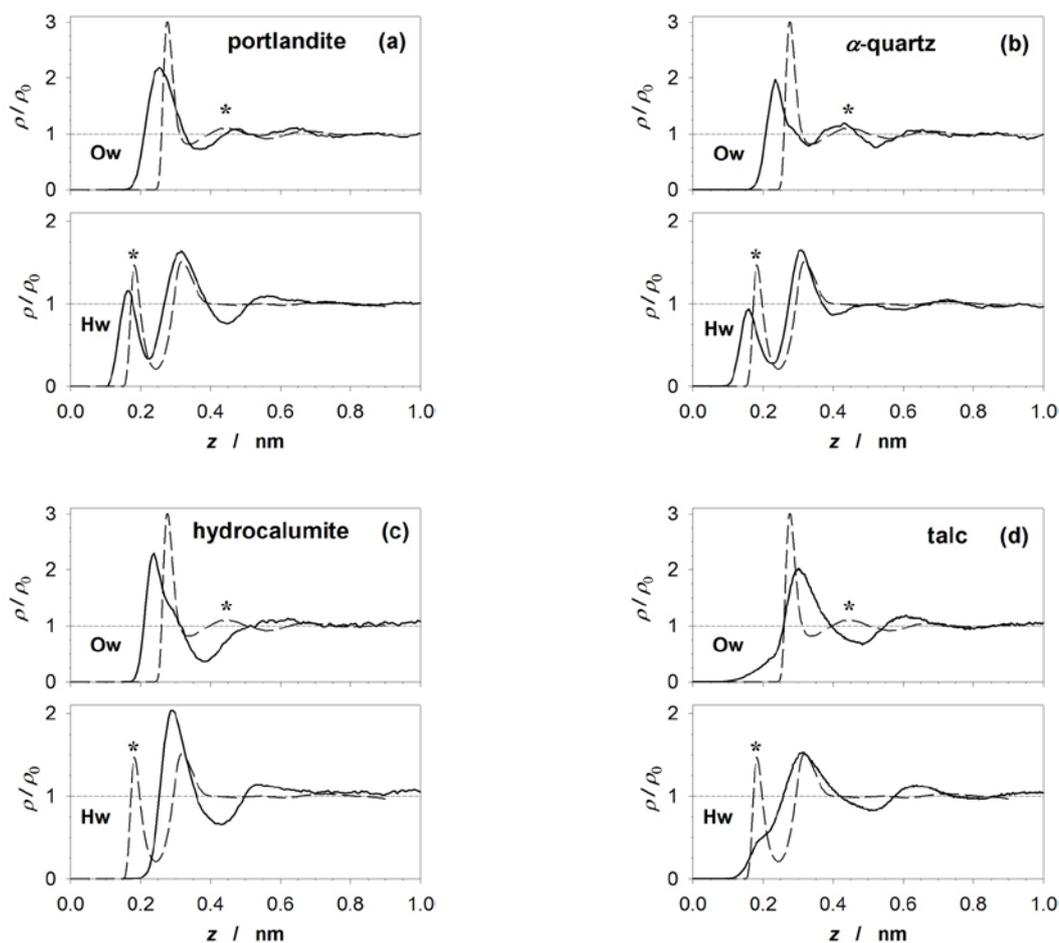


Figure 2: Contour maps of the atomic density of oxygens of H_2O (O_w) within the first molecular layer of water at the fully hydroxylated (001) surface of quartz, which is shown below the contours as tetrahedra (Si) and dark (O) and white (H) balls . See Fig. 1 for the detailed notations.

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This is clearly visible from the comparison of the atomic density profiles for O_w and H_w of these surfaces, which provide important additional insight into the ordering of water molecules at the interface. Two orientations of H_2O molecules are energetically most favorable for the formation of stable hydrogen bonds to the hydroxylated surfaces of portlandite and quartz (Fig.3a,b). In one of them, water molecules have one H-atom directed toward an O-atom of the surface hydroxyl, thus donating an H-bond to the surface. This is reflected in the presence of a strong peak at ~ 0.2 nm on the H_w density profiles. In this orientation, the second H-atom contributes to the peak at ~ 0.35 nm. However, a stronger contribution to the 0.35 nm peak in the H_w distribution arises from the other favorable H_2O orientation, in which it accepts an H-bond donated by the surface hydroxyl groups. In this case, both H_{H_2O} of the surface water molecule contribute to the density profile peak at 0.35 nm. In both preferred orientations of the surface water molecules, their O-atoms are located approximately at the same distance from the surface (~ 0.25 - 0.28 nm). The ability of the surface water molecules to both donate and accept H-bonds to the substrate surface creates a very well developed H-bonding network across the interface with the structure strongly resembling that of bulk liquid water. O-O and O-H radial distribution functions of liquid water are shown as thick dashed lines in Fig. 3 to illustrate this similarity. The $g_{OH}(r)$ peak at ~ 0.2 nm is a definitive signature of donated hydrogen bonds in the system, while the $g_{OO}(r)$ peak at ~ 0.45 nm clearly indicates that the local ordering of the H-bonding network is close to tetrahedral – a typical characteristic of bulk liquid water structure.



147 Figure 3: Density profiles of O_w and H_w atoms of the interfacial H_2O molecules as functions of the distance from the
 148 surface (z) of four different mineral substrates. Origin of z is defined here for all surfaces as the average Z-coordinate of
 149 the oxygen atoms at the substrate surface. The local density (ρ) is normalized with respect to the corresponding atomic
 150 density in bulk liquid water (ρ_0) and the radial distribution functions $g_{OO}(r)$ and $g_{OH}(r)$ of bulk liquid water are shown as
 151 thick dashed lines for comparison. Stars mark the characteristic H-bonding features of these functions.
 152

153 For hydrocalumite (Fig.3c) the computed near-surface water structure is strikingly different. Due to the
 154 positive structural charge of this substrate, the H-bond donation to the surface by H_2O molecules is prevented and the

155 well-interconnected H-bonding network is not formed in the interfacial region [11], despite the fact that the H₂O
 156 molecules are strongly attracted to this surface and it is clearly hydrophilic.

157 The atomic density profiles of the surface water on talc (Fig.3d) show that H₂O molecules are located
 158 somewhat farther away than from the previously discussed hydrophilic surfaces (Figs.3a-c) and the H₂O molecular
 159 orientation is predominately parallel to the surface (the peaks of O_w and H_w are approximately at the same distance of
 160 0.3-0.32 nm), in accordance with the hydrophobic character of talc. However, donation of weak H-bonds is still
 161 possible to this surface, as is clearly evident from the shoulders of the O_w and H_w distributions around 0.2 nm. Thus, talc
 162 surface exhibits a weak degree of hydrophilicity under ambient conditions, which can be most pronounced at very low
 163 relative humidity [33], or can even be turned into a strongly hydrophilic surface by the application of very high
 164 pressures and temperatures typical for the Earth's mantle conditions [40,41].

166 3.2. Interfacial water dynamics

167 The simulated diffusional dynamics of surface water is in agreement with these observations: the H₂O
 168 molecules have noticeably higher mobility at the hydrophobic surface of talc and significantly lower mobility at all
 169 hydrophilic surfaces [18,20]. Longer-time-scale MD simulations provide an additional opportunity to quantify these
 170 relatively slow diffusional motions of interfacial by means of the Van Hove self-correlation function (VHSCF) [9],
 171

$$172 \quad G_S(r, t) = \frac{1}{N} \left\langle \sum_{i=1}^N \delta[r + r_i(0) - r_i(t)] \right\rangle, \quad (1)$$

173 where r is the coordinate of the molecule, t is time, and N is the total number of molecules. This relationship describes
 174 the correlation in the positions of the same atom at different times, such that $4\pi r^2 G_S(r, t) dr$ is the probability of finding
 175 an atom at distance r after a time t if the position of this atom was at the origin $r = 0$ at the initial time $t = 0$. The Fourier
 176 transform of the VHSCF represents the incoherent or self-intermediate scattering function (SISF), which can be directly
 177 measured in incoherent quasielastic neutron scattering experiments and contains detailed information concerning the
 178 single-molecule dynamics both in time and space domain.
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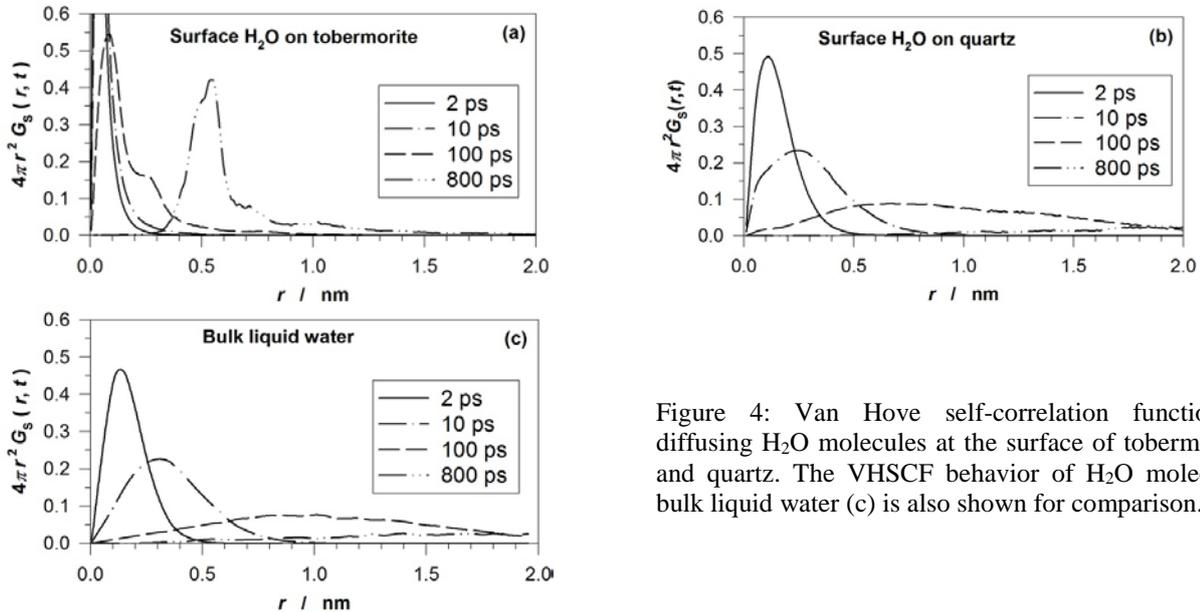


Figure 4: Van Hove self-correlation functions for diffusing H₂O molecules at the surface of tobermorite (a) and quartz. The VHSCF behavior of H₂O molecules in bulk liquid water (c) is also shown for comparison.

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182 The calculation of this self-correlation function for the water molecules strongly bound to the surface of
 183 tobermorite, which exhibits nanoscale surface roughness strongly affecting the first two monolayers of H₂O, shows a
 184 dramatic difference in the dynamic behavior of adsorbed H₂O (Fig. 4a), as compared with the diffusional dynamics in
 185 bulk liquid water (Fig. 4c), and yields a characteristic time scale of the diffusional process of the order of $\tau_m \sim 0.8$ ns, in
 186 excellent agreement with NMR measurements, indicating a mean time between diffusional jumps on the surface of
 187 ~ 0.8 - 1.0 ns [43]. The calculated long-time-scale VHSCF also points to a hopping diffusional mechanism with a length
 188 scale of $l_m \sim 5.5$ Å, i.e., close the characteristic lattice dimension of crystalline ice between two neighboring strong H-
 189 bonding sites. From the two-dimensional Einstein equation, $D = \langle |\mathbf{r}(t) - \mathbf{r}(0)|^2 \rangle / 4t$, the mean time for jumps between

190 surface sites from the NMR results yields a diffusion coefficient of $0.9 \times 10^{-10} \text{ m}^2/\text{s}$, which is also in remarkable
191 agreement with the average diffusion coefficient for all surface-associated H_2O molecules obtained from MD
192 simulations ($1.0 \times 10^{-10} \text{ m}^2/\text{s}$). Here, r is the mean jump displacement (assumed to be 5.5 \AA , from the results of the
193 VHSCF calculations, Fig.4a) and τ is the mean jump time, $\tau \sim \tau_m \sim 0.8 \text{ ns}$. This level of agreement definitely provides
194 strong support for the interpretation of the experimental results and strong encouragement for further application of
195 atomistic computational modeling techniques to study the molecular scale properties of nano-confined water.

196 The hopping surface diffusion mechanism is not evident for the water molecules at an atomically smooth
197 surface, such as that of quartz (Fig.4b): the corresponding VHSCFs show only slight differences with the diffusional
198 dynamics in bulk liquid water (Fig.4c).

200 3.3. Effect of metal cations on the membrane fouling

201 Prediction of ultrafiltration membrane performance is not generally possible without performing pilot-scale
202 tests because membrane fouling is related to membrane material and solution chemistry. The inadequate understanding
203 of membrane fouling has in many cases hindered wider adoption of membrane processes in large-scale drinking water
204 treatment plants. Computational molecular modeling of the model membrane-solution interfaces allowed detailed
205 probing of the fouling process due to the interaction of dissolved natural organic matter (NOM) with various metal
206 cations [39]. It was observed that divalent ions (Ca^{2+} and Mg^{2+}) may cause membrane fouling not by forming “ionic
207 bridges” between the negatively charged functional groups on the membrane surface and the negatively charged
208 functional groups of NOM, but rather by promoting the aggregation of NOM molecules in solution.

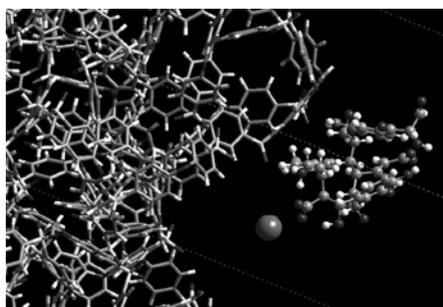


Figure 5: A fragment of the MD simulation box containing the polyethersulfone membrane surface (cylinder representation), a calcium ion (large dark ball), and a molecule of natural organic matter (ball and stick representation) coordinating to Ca^{2+} by its deprotonated carboxylic group. All interfacial water molecules are hidden for clarity.

210 Thus, even though the presence of truly dissolved NOM is not important in the fouling process, it can
211 contribute to fouling after being aggregated by divalent ions. The partially neutralized Ca^{2+} -NOM complex (Fig. 5) may
212 be more easily adsorbed at the membrane surface than a negatively charged NOM molecule. However, it still remains
213 unclear whether fouling is caused by the cation-mediated NOM aggregation in solution or by stronger NOM-surface
214 interactions in the presence of divalent ions. Since only one NOM molecule was used in the present model, NOM
215 aggregation could not be quantitatively studied in this work. More extensive MD simulations using multiple NOM
216 molecules, employing much larger size of the simulated systems, and much longer simulation times (at least on the
217 order of tens of nanoseconds) are necessary to quantitatively address these problems.

219 4. CONCLUSIONS

220 The methods of computational molecular modeling in application to the mineral-fluid systems have advanced
221 dramatically in the past decade and have already emerged as very powerful quantitative tools in the studies of the
222 structure and properties of interfacial and nano-confined aqueous solutions. There is a growing consensus, however,
223 that the idealized mineral-fluid interfaces typically used in the present-day molecular simulations and the relatively
224 small assemblages used in these models and in quantum calculations are at least an order of magnitude too small in
225 characteristic length (three orders of magnitude in volume) to effectively address many critically important issues. A
226 typical size of the systems currently simulated is on the order of a few nanometers, but can reach up to 20 nm [26] for
227 classical force-field-based MD simulations which can probe the systems evolution over tens of nanoseconds. However,
228 the application of quantum MD is still limited to the sizes $< 1.5\text{-}2 \text{ nm}$ and time periods $\sim 10 \text{ ps}$. The inability to
229 effectively model realistic compositional and structural complexity is significantly impeding further progress. For
230 instance, idealized, defect-free mineral surfaces are of little use in predicting local charge development, which is key to
231 understanding surface reactivity in real systems [13,16]. Particle edges, surface defects, the surface roughness on length
232 scales of 10 to 100 nm play central roles in determining the energetics and kinetics of dissolution, precipitation, sorption
233 and catalytic reactions at mineral interfaces, but these phenomena are currently modeled only very rarely and on a
234 limited scale [19,21,34,35]. There is no doubt, however, that further rapid development of supercomputing capabilities

235 will allow such molecular simulations to achieve truly realistic system sizes and time scales, thus significantly
236 increasing their practical value.

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243 REFERENCES

- 244 [1] Brown GE (2001) Surface science - How minerals react with water. *Science* 294:67-69
245 [2] Sverjensky DA (2006) Prediction of the speciation of alkaline earths adsorbed on mineral surfaces in salt solutions.
246 *Geochim Cosmochim Acta* 70:2427- 2453
247 [3] Brown GE, Calas G (2012) Mineral-aqueous solution interfaces and their impact on the environment. *Geochem*
248 *Perspectives* 1(4-5):483-742
249 [4] Fenter P, Sturchio NC (2004) Mineral-water interfacial structures revealed by synchrotron X-ray scattering. *Progr*
250 *Surf Sci* 77:171-258
251 [5] Wenk H-R (ed) (2006) Neutron Scattering in Earth Sciences. *Rev Mineral Geochem* 63:1-471
252 [6] Shen YR, Ostroverkhov V (2006) Sum-frequency vibrational spectroscopy on water interfaces: Polar orientation of
253 water molecules at interfaces. *Chem Rev* 106:1140-1154
254 [7] Cole DR, Mamontov E, Rother G (2009) Structure and dynamics of fluids in microporous and mesoporous earth and
255 engineered materials. In: Liang L, Rinaldi R, Schober H (eds) *Neutron Applications in Earth, Energy and*
256 *Environmental Sciences*. Springer, New York.
257 [8] Arbogast T (ed) (2007) Computational Needs for the Subsurface Sciences. Workshop Report. U.S. DOE Office of
258 Science, April 2007, 291pp
259 [9] Frenkel D, Smit B (2002) *Understanding Molecular Simulation: From Algorithms to Applications*. 2nd ed,
260 Academic Press, San Diego.
261 [10] Cygan RT, Kubicki JD, (eds) (2001) *Molecular Modeling Theory and Applications in the Geosciences*. *Rev*
262 *Mineral Geochem* 42:1-531
263 [11] Kalinichev AG, Kirkpatrick RJ (2002) Molecular dynamics modeling of chloride binding to the surfaces of Ca
264 hydroxide, hydrated Ca-aluminate and Ca-silicate phases. *Chem Mater* 14:3539-3549
265 [12] Cygan RT, Liang J-J, Kalinichev AG (2004) Molecular models of hydroxide, oxyhydroxide, and clay phases and
266 the development of a general force field. *J Phys Chem B* 108:1255-1266
267 [13] Rustad JR, Felmy AR (2005) The influence of edge sites on the development of surface charge on goethite
268 nanoparticles: A molecular dynamics study. *Geochim Cosmochim Acta* 69:1405-1411
269 [14] Tossell JA (2005) Theoretical study on the dimerization of Si(OH)₄ in aqueous solution and its dependence on
270 temperature and dielectric constant. *Geochim Cosmochim Acta* 69:283-291
271 [15] Criscenti LJ, Kubicki JD, Brantley SL (2006) Silicate glass and mineral dissolution: Calculated reaction paths and
272 activation energies for hydrolysis of a Q³ Si by H₃O⁺ using ab initio methods. *J Phys Chem A* 110:198-206
273 [16] Kerisit S, Rosso KM (2006). Computer simulation of electron transfer at hematite surfaces. *Geochim Cosmochim*
274 *Acta* 70:1888-1903
275 [17] Spagnoli D, Cooke DJ, Kerisit S, Parker SC (2006) Molecular dynamics simulations of the interaction between the
276 surfaces of polar solids and aqueous solutions. *J Mater Chem* 16:1997-2006
277 [18] Wang J, Kalinichev AG, Kirkpatrick RJ (2006) Effects of substrate structure and composition on the structure,
278 dynamics and energetics of water on mineral surfaces: MD modeling study. *Geochim Cosmochim Acta* 70:562-582
279 [19] Churakov SV (2007) Structure and dynamics of the water films confined between edges of pyrophyllite: A first
280 principle study. *Geochim Cosmochim Acta* 71:1130-1144
281 [20] Kalinichev AG, Wang J, Kirkpatrick RJ (2007) Molecular dynamics modeling of the structure, dynamics and
282 energetics of mineral-water interfaces: Application to cement materials. *Cem Concr Res* 37:337-347
283 [21] Larentzos JP, Greathouse JA, Cygan RT (2007) An ab initio and classical molecular dynamics investigation of the
284 structural and vibrational properties of talc and pyrophyllite. *J Phys Chem C* 111:12752-12759
285 [22] de Leeuw NH, Cooper TG (2007) Surface simulation studies of the hydration of white rust Fe(OH)₂, goethite α -
286 FeO(OH) and hematite α -Fe₂O₃. *Geochim Cosmochim Acta* 71:1655-1673

- 287 [23] Predota M, Cummings PT, Wesolowski DJ (2007) Electric double layer at the rutile (110) surface.
288 3. Inhomogeneous viscosity and diffusivity measurement by computer simulations. *J Phys Chem C* 111:3071-3079
- 289 [24] Allen JP, Gren W, Molinari M, Arrouvel C, Maglia F, Parker SC (2009) Atomistic modelling of adsorption and
290 segregation at inorganic solid interfaces. *Molecular Simulation* 35: 584-608
- 291 [25] Cygan RT, Greathouse JA, Heinz H, Kalinichev AG (2009) Molecular models and simulations of layered
292 materials. *J Mater Chem* 19:2470-2481
- 293 [26] Suter JL, Anderson RL, Greenwell HC, Coveney PV (2009) Recent advances in large-scale atomistic and coarse-
294 grained molecular dynamics simulation of clay minerals. *J Mater Chem* 19:2482-2493
- 295 [27] Wang J, Kalinichev AG, Kirkpatrick RJ (2009) Asymmetric hydrogen bonding and orientational ordering of water
296 at hydrophobic and hydrophilic surfaces: A comparison of water/vapor, water/talc, and water/mica interfaces.
297 *J Phys Chem C* 113:11077-11085
- 298 [28] Argyris D, Cole DR, Striolo A (2010) Ion-specific effects under confinement: The role of interfacial water. *ACS*
299 *Nano* 4:2035-2042
- 300 [29] Bourg IC, Sposito G (2010) Connecting the molecular scale to the continuum scale for diffusion processes in
301 smectite-rich porous media. *Env Sci Technol* 44:2085-2091
- 302 [30] Malikova N, Dubois E, Marry V, Rotenberg B, Turq, P (2010) Dynamics in clays - combining neutron scattering
303 and microscopic simulation. *Zeitschrift Phys Chem* 224:153-181
- 304 [31] Fenter P, Lee SS, Skelton AA, Cummings PT (2011) Direct and quantitative comparison of pixelated density
305 profiles with high-resolution X-ray reflectivity data. *J Synchrotron Radiat* 18:257-265
- 306 [32] Ferrage E, Sakharov BA, Michot LJ, Delville A, Bauer A, Lanson B, Grangeon S, Frapper G, Jimenez-Ruiz M,
307 Cuello GJ (2011) Hydration properties and interlayer organization of water and ions in synthetic Na-smectite with
308 tetrahedral layer charge. 2. Toward a precise coupling between molecular simulations and diffraction data.
309 *J Phys Chem C* 115:1867-1881
- 310 [33] Rotenberg B, Patel AJ, Chandler D (2011) Molecular explanation for why talc surfaces can be both hydrophilic
311 and hydrophobic. *J Amer Chem Soc* 133: 20521-20527
- 312 [34] Liu X, Lu X, Meijer EJ, Wang R, Zhou H (2012) Atomic-scale structures of interfaces between phyllosilicate
313 edges and water. *Geochim Cosmochim Acta* 81:56-68
- 314 [35] Tazi S, Rotenberg B, Salanne M, Sprik M, Sulpizi M (2012) Absolute acidity of clay edge sites from ab-initio
315 simulations. *Geochim Cosmochim Acta* 94:1-11
- 316 [36] Sahai N (2002) Is silica really an anomalous oxide? Surface acidity and aqueous hydrolysis revisited. *Env Sci*
317 *Technol* 36:445-452
- 318 [37] Morrow CP, Yazaydin AO, Krishnan M, Bowers GM, Kalinichev AG, Kirkpatrick RJ (2013) Structure, energetics,
319 and dynamics of smectite clay interlayer hydration: Molecular dynamics and metadynamics investigation of Na-
320 hectorite. *J Phys Chem C* 117:5172-5187
- 321 [38] Duval Y, Mielczarski JA, Pokrovsky OS, Mielczarski E, Ehrhardt JJ (2002) Evidence of the existence of three
322 types of species at the quartz-aqueous solution interface at pH 0-10: XPS surface group quantification and surface
323 complexation modeling. *J Phys Chem B* 106:2937-2945
- 324 [39] Ahn WY, Kalinichev AG, Clark MM (2008) Effects of background cations on the fouling of polyethersulfone
325 membranes by natural organic matter: Experimental and molecular modeling study. *J Membr Sci* 309:128-140
- 326 [40] Wang JW, Kalinichev AG, Kirkpatrick RJ (2005) Structure and decompression melting of a novel, high-pressure
327 nanoconfined 2-D ice. *J Phys Chem B* 109:14308-14313
- 328 [41] Wang JW, Kalinichev AG, Kirkpatrick RJ (2004) Molecular modeling of the 10Å phase at subduction zone
329 conditions. *Earth Planet Sci Lett* 222:517-527
- 330 [42] Kalinichev AG (2001) Molecular simulations of liquid and supercritical water: Thermodynamics, structure, and
331 hydrogen bonding. *Rev Mineral Geochem* 42:83-129
- 332 [43] Korb JP (2010) Multi-scales nuclear spin relaxation of liquids in porous media. *Compt Rend Phys* 11:192-203
- 333